Nuclear Quadrupole Moment of ^{57 m}Fe from α-Fe₂O₃ Data*

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The electric-field gradient at the Fe³⁺ site in hematite (α-Fe₂O₃) has been recomputed by use of the ionic model, including contributions of electronic dipole and quadrupole moments of the oxygen ion, as well as apparent quadrupolar effects due to thermal vibration of oxygen. The calculations have been performed self-consistently for various parameter sets of dipole and quadrupole polarizabilities, α_D and α_Q . As in corundum (α -Al₂O₃), the field gradient is insensitive to α_D because of the (accidentally) nearly vanishing field at the oxygen position. α_0 , however, is a sensitive parameter. The vibrational contribution is small, being \sim 5% of the total field gradient. Assuming a dipole polarizability $\alpha_D(\mathrm{O}^{2-})=2.91\,\mathrm{\AA}^3$, we find a range of values for the nuclear quadrupole moment of the 14.4-keV excited state of ^{57m}Fe, depending on $\alpha_Q(O^{2-})$, which is not well known. For $\alpha_Q(O^{2-}) = 0$, $Q(5^{7m}\text{Fe}) = 0.29$ b. For $\alpha_Q = 0.3$, Q = 0.16 b.

INTRODUCTION

NUMBER of attempts have been made to determine the nuclear electric quadrupole moment Q of the 14.4-keV excited state of ^{57m}Fe from nuclear quadrupole coupling constants e^2qQ obtained from Mössbauer-effect spectroscopy.1-7 One of the typical approaches has been to use simple, predominantly ionic crystal structures, such as α-Fe₂O₃ (hematite), and to calculate the electric-field gradient at the Fe³⁺ site by summation over point-charge or point-dipole lattices. The difficulty with this approach is that the point multipole expansion converges slowly, especially if it is used for the computation of ionic-potential derivatives at cationic sites in oxidic crystals. The reason is the critical dependence of the derivatives on the unknown ionic polarizabilities of oxygen. α-Fe₂O₃ is a good example of the inadequacy of the simple point-dipole approach.

In this article, self-consistent electric-field gradients at the Fe3+ site in hematite are calculated by use of the ionic model, including dipole and quadrupole moments of the oxygen ion. The calculations have been performed for various sets of polarizabilities to illustrate the dependence on these parameters. A full description of the calculation, including notation and units, may be found in a report by Hafner and Raymond⁸ on corundum (α -Al₂O₃), which is isomorphous to hematite.

A discrepancy in the value of Q based on data for ferric versus ferrous compounds has been noted. We investigate here the effect of including induced quadrupole moments on this discrepancy.

(1968).

SELF-CONSISTENT IONIC-POTENTIAL DERIVATIVES AT LATTICE SITES

Crystal fields in hematite are best described in the rhombohedral system which can be transformed to a rectangular system with three equivalent axes without loss of symmetry (cf. Table I). In this system, the ionic potentials and their first and second derivatives at the Fe3+ and O2- sites in hematite are completely determined by eight potential components V_i , which are related to each other by the equations

$$V_{i} = V_{i}^{M} + \sum_{j=3}^{8} K_{ij}\alpha_{j}V_{j}.$$
 (1)

 V_1 and V_2 are the Madelung potentials, $\sqrt{3}V_3$ $=V_r(\text{Fe}^{3+})$, and $\sqrt{2}V_5=V_S(\text{O}^{2-})$ are the potential gradients at Fe³⁺ and O²⁻ sites, respectively. V_4 $=V_{rr}(Fe^{3+})$ is the second derivative (negative field gradient) at Fe³⁺ parallel to the trigonal axis. V_6 , V_7 , and V_8 are the second derivatives (negative fieldgradient components) V_{xx} , V_{xy} , and V_{xz} at the oxygen sites. The dipole polarizability of Fe³⁺($-\alpha_3$) and the quadrupole polarizability of Fe³⁺ $(-\alpha_4)$ are assumed to be negligible. 9 $-\alpha_5 = \alpha_D$, and $-\alpha_6 = -\alpha_7 = -\alpha_8 = \alpha_Q$, respectively, represent the dipole and quadrupole polarizabilities of O²⁻. For more details about the notation the reader is referred to Ref. 8. The monopole terms V_{i}^{M} and the crystallographic K_{ij} factors have to be found by lattice summation. For this the Ewald-

Table I. Transformation matrix rhombohedral (a_r,α_r) -rectangular (x,y,z).

	$\begin{pmatrix} a & b \\ b & a \\ b & b \end{pmatrix}$		a = 5.01994 $b = 1.45985$	
Fe O	u,u,u; etc. $v, -v, 0$; etc.	u = 0.10528 a $v = 0.3072$ a	$a_h = 5.0346 \text{ Å}$ $c_h = 13.752 \text{ Å}$	

^a Blake, Finger, and Zoltai (Ref. 10).

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⁵ F. S. Ham, Phys. Rev. 160, 328 (1967).

⁶ M. Weissbluth and J. E. Maling, J. Chem. Phys. 47, 4166 (1967). (1967).

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⁹ The potential gradient V_{τ}^{M} at the Fe³⁺ site due to the point charges is 0.2712×10^{6} esu. The dipole polarizability α of Fe³⁺ is unknown. Assuming $\alpha(\text{Fe}^{3+}) = 0.1 \,\text{Å}^3$ and using K_{35} (Table II) we find a value of 0.0217 × 1014 esu as contribution to the field gradient $V_{rr}(\mathrm{Fe^{3+}})$.

TABLE II. Crystallographic Kij factors of hematite (cgs units).

j	3	4	5	6	7	8
3 4	-0.0766×10^{24}		$+0.5402\times10^{24}$			
5 6 7 8	+0.5402×10 ²⁴	$-0.2811 imes 10^{32} \ +0.4434 imes 10^{40} \ -0.1304 imes 10^{40} \ -0.2977 imes 10^{40}$	$+0.1277\times10^{24} \\ +0.0074\times10^{32} \\ +0.0895\times10^{32} \\ -0.0661\times10^{32}$	$+0.0098 imes 10^{32} \ +0.0424 imes 10^{40} \ +0.0105 imes 10^{40} \ -0.0477 imes 10^{40}$	$+0.3583\times10^{32} \\ +0.0316\times10^{40} \\ +0.0639\times10^{40} \\ -0.0023\times10^{40}$	$\begin{array}{c} -0.1323 \times 10^{32} \\ -0.0715 \times 10^{40} \\ -0.0011 \times 10^{40} \\ +0.0368 \times 10^{40} \end{array}$

a Definitions cf. Hafner and Raymond (Ref. 8).

potential sums were expanded in spherical harmonics. The summation was performed by use of the structural parameters of Blake, Finger, and Zoltai¹⁰ (Table I). Summations over about 1500 ions were found to be sufficient for obtaining four significant figures for the V_{i}^{M} terms. The K_{ij} factors were obtained by the pointcharge cluster technique. It should be noted that the K_{ij} 's are purely geometrical factors, and it is irrelevant whether the multipoles are of an ionic or nonionic nature. The factors are given in Table II. They can be compared with the previously reported values in Table III. They are rather similar to those of corundum, α-Al₂O₃ (Table III of Ref. 8). In contrast to the monopole terms they are not very sensitive to slight changes of atomic coordinates.8

RESULTS AND DISCUSSION

The ionic-field gradient at the Fe³⁺ position can be calculated for any combination of polarizability parameters by use of V_{i}^{M} , K_{ij} , and solution of the system of partially simultaneous equations (1). As with corundum, the polarizabilities are not well known. In the previous calculations, the dipole polarizability $\alpha_D = 2.91 \text{ Å}^3$, derived from the index of refraction by

Table III. Comparison with previous data of point-charge potential components and K factors.

Notation ^a	Units	This work	AMW b
$V_{r^{M}}(\mathrm{Fe})$	10 ⁶ esu	-0.27120	-0.27469
$V_{s}^{M}(O)$	$10^6 \mathrm{esu}$	0.06228	0.06589
$V_{rr}^{M}(\mathrm{Fe})$	1014 esu	0.4261	0.42422
$V_{xx}^{M}(O)$	1014 esu	-1.3643	
$V_{xy}^{M}(O)$	1014 esu	-1.8006	
$V_{xz}^{M}(O)$	1014 esu	0.4879	
$K_1 = -K_{33}$	$10^{24}~{\rm cm^{-3}}$	0.0766	0.09032
$K_2^{\text{c,d}} = -(\sqrt{3}/\sqrt{2})K_{35}^{\text{q}}$	$10^{24}~{\rm cm}^{-3}$	-0.6616	-0.66161 d
$K_4^{\text{c}} = -K_{55}$	$10^{24}~{\rm cm}^{-3}$	-0.1277	-0.03358 $^{\circ}$
eq_{d-0} °	10^{32} esu	-0.3975	-0.39736

Reference 8.

the Clausius-Mossotti relation, was generally used. Here, we have calculated the field gradient for this polarizability, and for $\alpha_D = 2.0 \text{ Å}^3$, as well as a number of different quadrupole polarizabilities α_Q . The total field gradient is plotted in Fig. 1 as a function of α_Q for both α_D values. It is evident that the field gradient is not sensitive to α_D , whereas it depends critically on α_0 . This is due to the fact that the electrostatic field at the oxygen position is accidentally small, while the field gradient at this position is large. With the quadrupole coupling constant¹ $e^2qQ = +0.880$ mm/sec and the antishielding factor¹¹ $\gamma_{\infty} = -9.14$, we find the quadrupole moments given in Table IV. The ionic multipole moments and the various contributions to the field gradient of Fe3+ are given in Table V. Thus, we find that a determination of Q by lattice-summation methods in Fe₂O₃ is possible only if $\alpha_Q(O^{2-})$ is accurately known. This fact has not been sufficiently recognized in previous work on this problem and therefore inadequate conclusions have been reached.

A few theoretical calculations of α_Q have been performed, using Hartree-Fock wave functions. The values obtained are 400-1000 Å⁵. 12,13 Using contracted

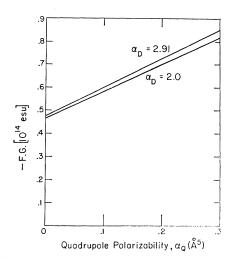


Fig. 1. Dependence of second derivative of the potential (negative electric-field gradient) at Fe³⁺ site in hematite on quadrupole polarizability of the oxygen ion.

b Reference 1.

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This inconsistency also exists in J. O. Artman and J. C. Murphy, Phys. Rev. 135, 1622 (1966) and in J. O. Artman, ibid. 143, 541 (1966).

¹⁰ We are grateful to Dr. R. L. Blake for providing us with the data prior to publication; coordinates and lattice constants are also quoted in Ref. 1.

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Table IV. Quadrupole moment of 57m Fe for various oxygen-ion polarizabilities.

$(\mathring{\mathbb{A}}^3)$	$egin{pmatrix} lpha_Q \ (\mathring{ m A}^5) \end{pmatrix}$	$-\text{F.G.} = V_{rr}(\text{Fe}^{3+})$ (×10 ¹⁴ esu)	$Q({}^{57m}{ m Fe})^{-{f a}}$ (b)
2.0	0	0.4655	+0.299
2.0	0.10	0.5858	0.237
2.0	0.20	0.7036	0.197
2.0	0.30	0.8189	0.170
2.0	0.50	1.0426	0.133
2.91	0	0.4786	0.290
2.91	0.10	0.6051	0.230
2.91	0.20	0.7292	0.191
2.91	0.30	0.8508	0.163
2.91	0.50	1.0920	0.127

a Using γ_{∞} (Fe³⁺) = -9.14.

wave functions, Burns and Wikner¹⁴ calculated 7 Å⁵. All these values are much too large. For instance, if we use $\alpha_Q = 5 \text{ Å}^5$ in our calculation, we find Q = 0.03 b. Lattice-sum calculations in corundum, 8 using a variable parameter approach, indicate an α_Q of about 0.2 Å⁵. Similar calculations in kyanite (triclinic Al₂SiO₅) ¹⁵ reckon $\alpha_Q(O^{2-})$ to be about 0.1 Å⁵. The polarizability of course depends on the crystal structure, so the range of values 0-0.5 Å⁵ was chosen for this calculation. Without more definite information on this polarizability however, the O(57mFe) remains quite uncertain.

It is to be noted that in corundum-type structures, the quadrupolar contribution to the field gradient at the cationic site appears to be larger than the dipolar contribution. This is due to the fact that the field at the oxygen site is (accidentally) nearly zero. It is slightly positive in corundum and slightly negative in hematite. The field gradient at O²-, however, is large. In crystal structures with more general point symmetries at the ionic sites (e.g., kyanite) the quadrupolar effects are found to be considerably smaller than the dipolar effects.

Table V. O²⁻ multipole moments and Fe³⁺ field gradients in hematite ($\alpha_D = 2.91 \text{ Å}^3$; $\alpha_Q = 0.1 \text{ Å}^5$).

O2-	Dipole moment	-0.1942×10^{-18}
O2-	Quadrupole momenta,b	
	\hat{Q}_{xx}	$+0.1354\times10^{-26}$
	Q_{xy}	$+0.1834\times10^{-26}$
	Q_{xz}	$-0.0494{ imes}10^{-26}$
Fe^{3+}	Field gradient (second	derivative of the potential)b
	Point charge	$+0.4261 \times 10^{14} \text{esu}$
	Dipole	$+0.0772\times10^{14} \text{ esu}$
	Quadrupole	$+0.1018 \times 10^{14} \text{ esu}$
	Total	$+0.6051 \times 10^{14} \text{ esu}$

^a x, y, z according to Table I. ^b Static contribution only.

TABLE VI. Thermal and electrostatically-induced quadrupole moments of the oxygen ion in α -Fe₂O₃.

8.	Thermal natural ^b 10 ⁻²⁶ esu	Thermal synthetic ^c 10 ⁻²⁶ esu	Electrostatic $\alpha_D = 2.91 \text{ Å}^3; \alpha_Q = 0.1 \text{ Å}^5$ 10^{-26} esu
Qxx Qxy Qxz Qyz Qyz Qzz	$\begin{array}{c} +0.012 \\ +0.019 \\ -0.044 \\ +0.012 \\ -0.044 \\ -0.024 \end{array}$	+0.008 $+0.038$ -0.021 $+0.008$ -0.021 -0.016	+0.135 $+0.183$ -0.049 $+0.135$ -0.049 -0.271

The anisotropic temperature coefficients given by Blake et al. 10,16 were used to calculate the thermal quadrupole moments at the oxygen site. The vibrational amplitudes were derived according to the formulation of Cruickshank.¹⁷ The thermal quadrupole moment is defined as $Q_{ij}=2e\langle u_{ij}^2\rangle$, where $\langle u_{ij}^2\rangle$ is the mean square amplitude of the thermal ellipsoid, 16 and e is the charge on the ion. The resulting moments are given in Table VI, with induced quadrupole moments given for comparison. The field gradient at the Fe³⁺ site due to thermal quadrupoles is $+0.033\times10^{14}$ esu for natural hematite and 0.0097×10¹⁴ esu for synthetic α-Fe₂O₃. The thermal vibrations of the oxygen ion at room temperature contribute 2-5% to the field gradient at Fe3+. This result is principally in agreement with that of Schwarzenbach¹⁸ who considered vibrational contributions to the field gradient at the Al3+ position in AlPO₄.

It has been pointed out that there is a disagreement between values of Q (57m Fe) obtained from ferric- and ferrous-compound data. Most recent calculations based on ferrous-compound data indicate a Q (57mFe) of about 0.2 b.^{3-5} Determinations of Q by other methods also result in Q of ~ 0.2 b.^{2,6} Van der Woude and Dekker calculated Q = +0.28 b in α -FeOOH (ferric), but they used only the monopole contribution. We have found that the O cannot be precisely determined in Fe₂O₃. Including the uncertainty due to the ionic polarization of O²⁻, and neglecting covalency, Q varies between 0.29 and 0.16 b.

It would be interesting to perform similar calculations for other ferric compounds. Of course, covalent participation in the predominantly ionic Fe³⁺-O²⁻ bonds will affect the external field gradient and Sternheimer factor at the Fe³⁺ site. But as long as the predominant, ionic contribution is still uncertain, due to insufficiently known O2- polarizabilities, estimates of covalent contributions are an involved problem.

¹⁴ G. Burns and E. G. Wikner, Phys. Rev. 121, 155 (1961). 15 M. Raymond, Ph.D. thesis, submitted to the University of Chicago (unpublished).

a x, y, z system according to Table I. b Reference 16. c Reference 10; the temperature coefficients of oxygen are (in units of 10*) 1=2.7, $\beta_{22}=5.4$, $\beta_{33}=0.2$, $\beta_{12}=2.7$, $\beta_{13}=0.25$, $\beta_{23}=0.5$. For dimensions cf. Ref. 16.

¹⁶ R. L. Blake, R. E. Hessevick, T. Zoltzi, and L. W. Finger, Am. Mineralogist **51**, 123 (1966).

¹⁷ D. W. J. Cruickshank, Acta Cryst. **9**, 747 (1956).

¹⁸ D. Schwarzenbach, Z. Kristallogr. **123**, 422 (1966).